# Universal Inversion: a Framework for Infusing Expert Knowledge in Bayesian Inverse Problems

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## February 28, 2023

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# Section 1

## Introduction: Inverse Problems and Stromboli Volcano

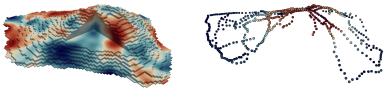
## Motivating Example: Stromboli Volcano

Want to learn the interior structure of the Stromboli volcano.

• Only allowed to measure gravitational field on the surface

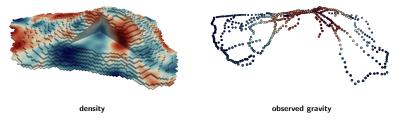


## Want to recover interior density field $\rho:D\to \mathbb{R}$ from surface gravity.



density

observed gravity



Properties:

- Linear operator data.
- "Large-scale": large 3 dimensional inversion grid.
- Sequential assimilation of new data and optimal design of experiment important in practice.

Gravity field  $G(s_i, \rho)$  at site  $s_i$  generated by underground density  $\rho$  can be written as linear operator:

$$y_i = G(s_i, \rho) = \int_D \rho(x)g(x, s_i)dx$$

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- Traditionally solved by discretizing on a grid  $\boldsymbol{x} = (x_1, ..., x_m)$ .
- Observation model

$$\boldsymbol{y} = \boldsymbol{G}(\rho(w_1), ..., \rho(w_r))^T =: \boldsymbol{G}\rho_{\boldsymbol{W}}$$

• Discretized observation operator  $\boldsymbol{G} \in \mathbb{R}^{n imes r}$  for n observations.

.

Assume  $\rho$  is a realization of a GP prior  $Z \sim GP(m, k)$ .

• Data model is  $Y = GZ_W + \epsilon$ .

• Approximate  $\rho$  by posterior, conditional on the observed data  $m{Y}=m{y}.$ 

- Posterior is Gaussian, fully characterized by:
  - mean vector  $\tilde{\boldsymbol{m}} = (\tilde{m}_{x_1},...,\tilde{m}_{x_m})^T$

• covariance matrix 
$$ilde{m{K}} = \left( ilde{k}(x_i,x_j)
ight)_{i,j=1,...,m}$$

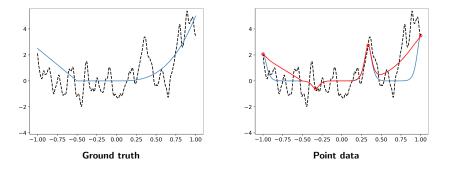
$$egin{aligned} ilde{m{m}} &= m{m} + m{K}m{G}^T \left(m{G}m{K}m{G}^T + m{\Delta}
ight)^{-1} \left(m{y} - m{G}m{m}
ight) \ ilde{m{K}} &= m{K} - m{K}m{G}^T \left(m{G}m{K}m{G}^T + m{\Delta}
ight)^{-1}m{G}m{K} \end{aligned}$$

Hard to implement in practice [TGL21]

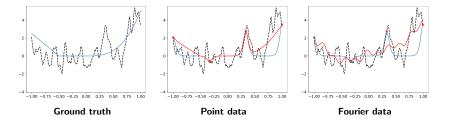
# Section 2

## Bayesian Inversion with Trends

## GP regression with trends known as universal kriging.



#### GP regression with trends known as **universal kriging**.



Goal: Extend to inverse problems to get "universal inversion".

## Trends allow inclusion of expert knowledge.

Model known geological stuctures such as:

- Depth-dependence: resulting from way volcanoes aggregate mass.
- Layers
- Chimneys: higher densities in central lava conduit.
- Fault lines: some volcanoes separate along fault through which high density magma infiltrates.

Trends modeled by specifying basis functions:

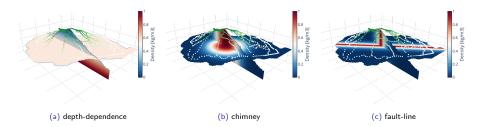


Figure: Trend basis functions used in the experiments. Solid balls denote the locations of the gravimetry observations (Stromboli dataset).

Assume prior is sum of trend + fluctuations described by Gaussian process

$$Z_x = \sum_{i=1}^p \beta_i f_i(x) + \eta_x,$$

- $\eta$  is a (centred) GP with kernel k,
- basis functions  $f_i: D \to \mathbb{R}$ ,
- trend coefficients  $\boldsymbol{\beta} = (\beta_1, ..., \beta_p)^T$ .

Put a Gaussian prior on the trend coefficients

$$\boldsymbol{\beta} \sim \mathcal{N}(\boldsymbol{\mu}, \boldsymbol{\Sigma})$$
.

#### Theorem

Conditionally on linear operator data  $Y = GZ_w + \epsilon$ , the posterior of the trend coefficients is Gaussian with mean and covariance given by:

$$\mathbb{E}\left[\boldsymbol{\beta}|\boldsymbol{y}\right] = \boldsymbol{\mu} + \Sigma F_{\boldsymbol{W}}^{T} G^{T} Q_{\boldsymbol{y}}^{-1} \left(\boldsymbol{y} - G F_{\boldsymbol{W}} \boldsymbol{\mu}\right)$$
$$\operatorname{Cov}\left(\boldsymbol{\beta}, \boldsymbol{\beta}|\boldsymbol{y}\right) = \Sigma - \Sigma F_{\boldsymbol{W}}^{T} G^{T} Q_{\boldsymbol{y}}^{-1} G F_{\boldsymbol{W}} \Sigma,$$

assuming that the matrix  $Q_{\boldsymbol{y}} := G \left( F_{\boldsymbol{W}} \Sigma F_{\boldsymbol{W}}^T + K_{\boldsymbol{W} \boldsymbol{W}} \right) G^T + \sigma_{\epsilon}^2 I$  is invertible.

Conditionally on the data, the distribution of Z is also that of a GP, with mean and covariance function given by:

$$m_{Z|\boldsymbol{y}}(\boldsymbol{x}) = F_{\boldsymbol{x}}\boldsymbol{\mu} + \left(F_{\boldsymbol{x}}\Sigma F_{\boldsymbol{W}}^{T} + K_{\boldsymbol{x}\boldsymbol{W}}\right)G^{T}Q_{\boldsymbol{y}}^{-1}\left(\boldsymbol{y} - GF_{\boldsymbol{W}}\boldsymbol{\mu}\right)$$
$$k_{Z|\boldsymbol{y}}\left(\boldsymbol{x}, \boldsymbol{x}'\right) = K_{\boldsymbol{x}\boldsymbol{x}'} + F_{\boldsymbol{x}}\Sigma F_{\boldsymbol{x}'}^{T} - \left(F_{\boldsymbol{x}}\Sigma F_{\boldsymbol{W}}^{T} + K_{\boldsymbol{x}\boldsymbol{W}}\right)G^{T}Q_{\boldsymbol{y}}^{-1}G$$
$$\left(F_{\boldsymbol{W}}\Sigma F_{\boldsymbol{x}'}^{T} + K_{\boldsymbol{W}\boldsymbol{x}'}\right)$$

## Volcano Inversion with Trend

Inversion with trends gives consistent results (data fit within noise std).

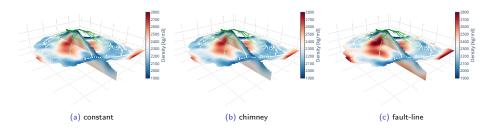


Figure: Posterior mean for the various trend models (Stromboli data).

But gives rise to new questions.

# Universal Inversion brings new challenges:

## Questions

- O Can we compare the performance of different trend models?
- Output Weight and Strength a
- Oan we perform measurements to "optimally" discriminate models?

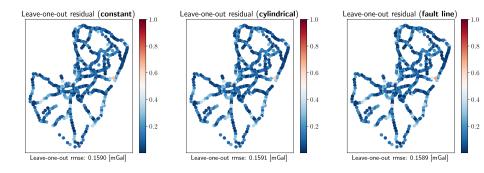
#### Idea

Given a data model  $Y = GZ_W$  and observed value y:

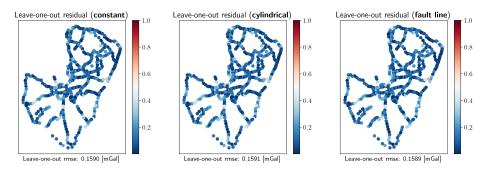
- **1** Remove one data point  $y_i$ .
- 2 Fit model on remaining data  $y^{(-i)}$ .
- **③** Compute mean of fitted model  $\hat{Z}^{(-i)} := \mathbb{E}[Z| \boldsymbol{Y}^{(-i)} = \boldsymbol{y}^{(-i)}]$
- Predict missing data using mean, compute residual  $e_i := y_i G\hat{Z}_{W}^{(-i)}$ .

 $\sum_{i=1}^{n} || \boldsymbol{y}_i - G\hat{Z}_{\boldsymbol{W}}^{(-i)} ||^2$  gives measure of model generalisation error.

# Leave-One Out (LOO) Cross-Validation on Stromboli



# Leave-One Out (LOO) Cross-Validation on Stromboli



# $\begin{array}{rl} \mbox{Gravimetric observations highly correlated} \\ \implies \mbox{LOO not informative.} \end{array}$

## Beyond LOO: Leave-k-out

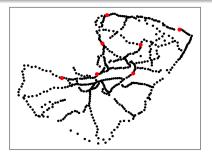
Can also consider subsets of size k: separate data in two batches:

$$\boldsymbol{y} = (\boldsymbol{y}_{\boldsymbol{i}}, \boldsymbol{y}_{-\boldsymbol{i}}),$$

where  $y_i$  has size k.

Want to compute residual when we predict  $y_i$  using  $y_{-i}$ 

$$e_i := y_i - G\hat{Z}_W^{(-i)}$$



## Problems

- Have to recompute full posterior at each CV pass.
- Full leave-k-out infeasible (combinatorial explosion).

- Need to select "representative" subsets.
- Need efficient formulae for computing residuals.

By generalizing [GS] all the information we need is contained in the augmented matrix:

$$\tilde{K} = \begin{pmatrix} GK_{WW}G^T & GF_W \\ F_W^TG^T & \mathbf{0} \end{pmatrix}.$$

We partition it as:

$$\tilde{K} = \begin{pmatrix} \tilde{K}_{ii} & \tilde{K}_{i-i} \\ \tilde{K}_{-ii} & \tilde{K}_{-i-i} \end{pmatrix} = \begin{pmatrix} G_{i \bullet} K_{WW} G_{\bullet i}^T & G_{i \bullet} K_{WW} G_{\bullet -i}^T & G_{i \bullet} F_W \\ G_{-i \bullet} K_{WW} G_{\bullet i}^T & G_{-i \bullet} K_{WW} G_{\bullet -i}^T & G_{-i \bullet} F_W \\ F_W^T G_{\bullet i}^T & F_W^T G_{\bullet -i} & \mathbf{0} \end{pmatrix}$$

where  $\tilde{K}_{ii} := G_{i \bullet} K_{WW} G_{\bullet i}^T$ .

CV residuals can the be computed by extracting subblocks of the inverse.

Residuals can be computed by extracting and inverting upper left block of the inverse:

$$\boldsymbol{e_i} = \left(\tilde{K}_{\boldsymbol{i}\boldsymbol{i}}^{-1}\right)^{-1} \left(\tilde{K}^{-1}\left[:,1:n\right]\boldsymbol{y}\right)_{\boldsymbol{i}},$$

where  $\tilde{K}_{ii}^{-1}$  denotes the ii sub-block of  $\tilde{K}^{-1}$  and y has dimension n.

Drasting computational savings by avoiding recomputation of full posterior.

## Application: k-fold CV to detect trends in Stromboli

Instead of considering all data subsets of a given size, consider k different subsets (folds).

• Remove one fold and predict using remaining ones.

**Choosing Folds** 

How can we choose the k subsets?

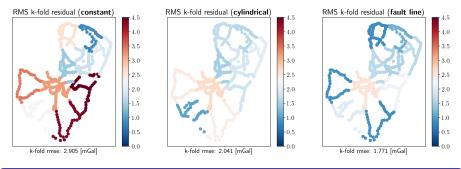
• Start with spatial clustering (heuristic).



Figure: Folds from kMeans clustering, k = 10.

## K-fold Cross-Validation: Stromboli (experimental)

## kMean CV favors fault line model.



## Question?

Is this representative of generalisation error?

Can consider the residuals as random variables

$$\boldsymbol{E_i} := \boldsymbol{G_i} \boldsymbol{Z_W} - \boldsymbol{G_i} \boldsymbol{\hat{Z}_W}^{(-i)},$$

where  $\hat{Z}^{(-i)}$  is the UK predictor based on evaluation of  $G_{-i}Z$ .

## Theorem

(under the model) The residuals are jointly Gaussian distributed, centred and with covariance

$$\operatorname{Cov}\left(E_{\boldsymbol{i}}, E_{\boldsymbol{j}}\right) = \left(\tilde{K}_{\boldsymbol{i}\boldsymbol{i}}^{-1}\right)^{-1} \tilde{K}_{\boldsymbol{i}\boldsymbol{j}}^{-1} \left(\tilde{K}_{\boldsymbol{j}\boldsymbol{j}}^{-1}\right)^{-1}$$

How can we leverage knowledge of theoretical distribution of residuals?

Choose folds to mimick generalisation error?

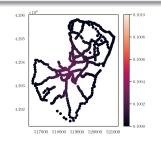


Figure: Theoretical CV variances (constant trend).

Gaussian hypothesis too "smooth" (few gravimetric observations kill all variance).



David Ginsbourger and Cedric Schärer, *Fast calculation of gaussian process multiple-fold cross-validation residuals and their covariances*.

Cédric Travelletti, David Ginsbourger, and Niklas Linde, *Uncertainty quantification* and experimental design for large-scale linear inverse problems under gaussian process priors, arXiv 2109.03457 (to appear in SIAM JUQ), 2021.